

# Radiative Transfer of Seismic Waves

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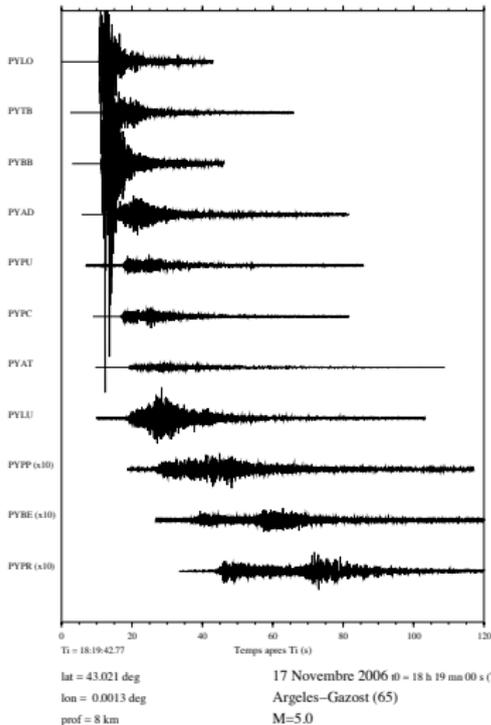
Atelier sur les ondes élastiques, Col de Porte, 13 janvier 2009

En collaboration avec: N. Le Bihan, M. Campillo, E. Larose, G. Nolet, C.  
Sens-Schönfelder, N. Shapiro, B. van Tiggelen, ...

# Content

- 1 Introduction to Radiative Transfer
- 2 Solutions of the Radiative Transfer Equation
- 3 Applications of Radiative Transfer

# Radiative Transfer in Seismology: Why?



## Observations

- Duration of signal  $\gg$  travel time of ballistic waves
- Rapid attenuation of direct waves
- “Cigar”-shaped envelope
- Pronounced **Coda**

## Role of Scattering

Crustal earthquake in the Pyrenees

# What is Radiative Transfer?

## Definition

Radiative Transfer is a theory aimed at predicting the spatio-temporal distribution of **energy** in a scattering medium

## The physical concepts of Radiative Transfer

- Local Energy Balance
- Angularly resolved energy flux: Specific intensity
- Scattering strength: mean free path

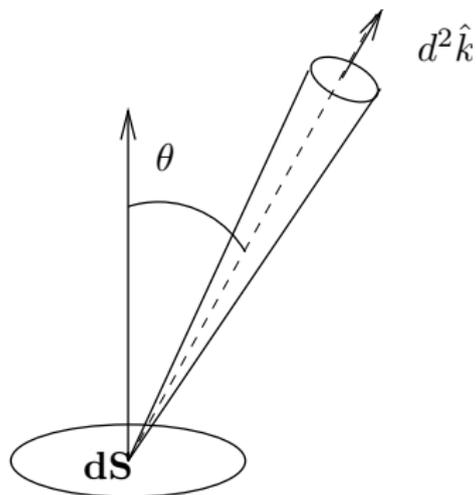
## Some references in seismology

- Introduced by Wu (1985) and developed by Aki, Zeng, Sato
- Monte Carlo simulations: Gusev and Abubakirov, Hoshiaba
- Sato and Fehler (Wave propagation and scattering in the heterogeneous Earth, Academic Press, 1998)
- Dmowska, Sato and Fehler, Eds, Vol. 50 of Advances in Geophysics, Academic Press, 2008
- IASPEI Task group: scattering and heterogeneity in the earth:  
<http://www.scat.geophys.tohoku.ac.jp/index.html>

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## The specific intensity



Angularly-resolved energy flux  
through a surface

### Definition

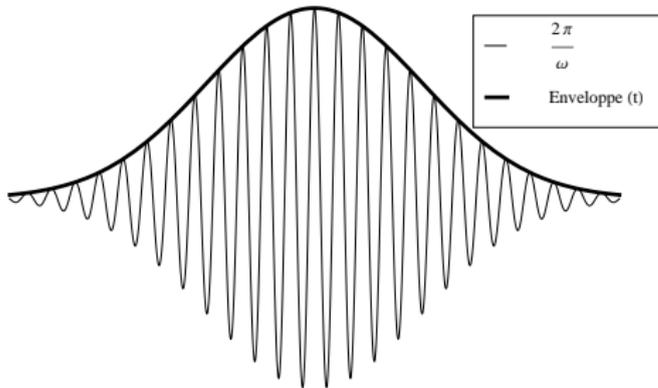
$\mathcal{I}(\omega, t, \mathbf{r}, \hat{\mathbf{k}}) \times dS \times \cos(\theta) \times dt \times d\omega =$  Amount of energy within the frequency band  $[\omega, \omega + d\omega]$  flowing through  $dS$  around direction  $\hat{\mathbf{k}}$  during time  $dt$

# Relation to the wavefield

## Wigner-Ville Spectrum

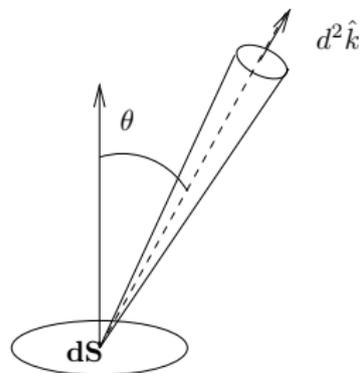
Consider a time-dependent field  $\psi(t)$ :

$$W_{\psi}(t, \omega) = \int_{-\infty}^{+\infty} \langle \psi(t - \tau/2) \psi(t + \tau/2)^* \rangle e^{-i\omega\tau} d\tau$$



- $\langle \cdot \rangle$ : Ensemble Average
- Separation of Time Scales
- $W_{\psi}$ : Instantaneous Energy Spectrum

# Wave content of the specific intensity



Angular Energy  
Spectrum

## Spatial Wigner-Ville

$$W_{\psi}(\mathbf{x}, \mathbf{k}) = \iiint \langle \psi(\mathbf{x} - \mathbf{r}/2) \psi(\mathbf{x} + \mathbf{r}/2)^* \rangle e^{i\mathbf{k} \cdot \mathbf{r}} d^3 r$$

- Slowly-modulated wave packet
- Specific intensity contains information on the **correlation** properties of the wavefield

# The Equation of Radiative Transfer

## Scalar Case

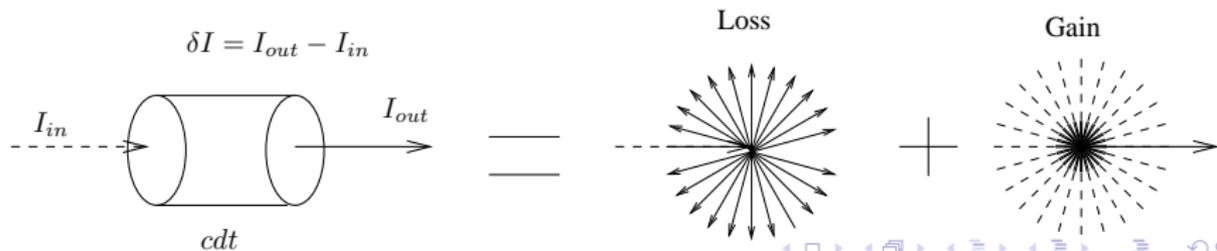
$$\left( \frac{1}{c} \frac{\partial}{\partial t} + \hat{\mathbf{k}} \cdot \nabla_{\mathbf{x}} \right) \mathcal{I}(t, \mathbf{x}, \hat{\mathbf{k}}) = - \left( \frac{1}{l^s} + \frac{1}{l^a} \right) \mathcal{I}(t, \mathbf{x}, \hat{\mathbf{k}}) + \frac{1}{l^s} \oint p(\hat{\mathbf{k}}, \hat{\mathbf{k}}') \mathcal{I}(t, \mathbf{x}, \hat{\mathbf{k}}') d^2 \hat{\mathbf{k}}'$$

$l^s$ : scattering mean free path

$l^a$ : absorption length

$c$ : wave speed

$p(\hat{\mathbf{k}}, \hat{\mathbf{k}}')$ : scattering anisotropy



# Assumptions in radiative transfer

## Averaging the products of 2 Green functions

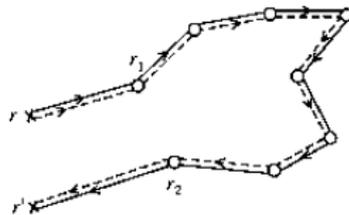
Example of scattering path



Visit the same scatterers



Pairing of trajectories



From Akkermans & Montambaux (2005)

## The case of vector waves

### The correlation tensor

$$\Gamma_{ij}(\mathbf{x}, \mathbf{k}) = \iiint \langle \psi_i(\mathbf{x} - \mathbf{r}/2) \psi_j(\mathbf{x} + \mathbf{r}/2)^* \rangle e^{i\mathbf{k} \cdot \mathbf{r}} d^3 r$$

Rotate from the frame  $(\hat{\mathbf{x}}_1, \hat{\mathbf{x}}_2, \hat{\mathbf{x}}_3)$  onto the frame  $(\hat{\mathbf{k}}_1 = \hat{\mathbf{k}}, \hat{\mathbf{k}}_2, \hat{\mathbf{k}}_3)$ :

$$\Gamma(\mathbf{x}, \mathbf{k}) = \begin{pmatrix} \Gamma_{\hat{\mathbf{k}}_1 \hat{\mathbf{k}}_1} & \Gamma_{\hat{\mathbf{k}}_1 \hat{\mathbf{k}}_2} & \Gamma_{\hat{\mathbf{k}}_1 \hat{\mathbf{k}}_3} \\ \Gamma_{\hat{\mathbf{k}}_2 \hat{\mathbf{k}}_1} & \Gamma_{\hat{\mathbf{k}}_2 \hat{\mathbf{k}}_2} & \Gamma_{\hat{\mathbf{k}}_2 \hat{\mathbf{k}}_3} \\ \Gamma_{\hat{\mathbf{k}}_3 \hat{\mathbf{k}}_1} & \Gamma_{\hat{\mathbf{k}}_3 \hat{\mathbf{k}}_2} & \Gamma_{\hat{\mathbf{k}}_3 \hat{\mathbf{k}}_3} \end{pmatrix}$$

### Interpretation

$\Gamma_{\hat{\mathbf{k}}_1 \hat{\mathbf{k}}_1}$ : *P*-wave intensity

$\Gamma_{\hat{\mathbf{k}}_2 \hat{\mathbf{k}}_2}, \Gamma_{\hat{\mathbf{k}}_3 \hat{\mathbf{k}}_3}$ : *S*-wave intensity along  $\hat{\mathbf{k}}_2, \hat{\mathbf{k}}_3$

$\Gamma_{\hat{\mathbf{k}}_2 \hat{\mathbf{k}}_3}, \Gamma_{\hat{\mathbf{k}}_3 \hat{\mathbf{k}}_2}$ : Stokes-like parameters  $\rightarrow$  polarization

**Remaining terms:** correlation between *P* and *S* waves

# The equation of radiative transfer for elastic waves

$$\left( \mathbf{c}^{-1} \frac{\partial}{\partial t} + \hat{\mathbf{k}} \cdot \nabla_{\mathbf{x}} \right) \mathcal{I}(t, \mathbf{x}, \hat{\mathbf{k}}) =$$

$$- \left( \mathbf{I}^{-1} + \mathbf{I}^{\mathbf{a}-1} \right) \mathcal{I}(t, \mathbf{x}, \hat{\mathbf{k}}) + \mathbf{I}^{-1} \int \mathbf{p}(\hat{\mathbf{k}}, \hat{\mathbf{k}}') \mathcal{I}(t, \mathbf{x}, \hat{\mathbf{k}}') d^2 \hat{\mathbf{k}}'$$

$$\mathbf{I} = \text{Diag}(I_p, I_s, I_s, I_s, I_s)$$

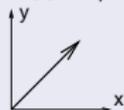
$$\mathbf{I}^{\mathbf{a}} = \text{Diag}(I_p^{\mathbf{a}}, I_s^{\mathbf{a}}, I_s^{\mathbf{a}}, I_s^{\mathbf{a}}, I_s^{\mathbf{a}})$$

$$\mathbf{c} = \text{Diag}(v_p, v_s, v_s, v_s, v_s)$$

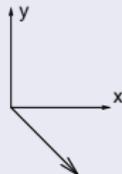
$$\mathcal{I} = (I_p, I_x, I_y, U, V)$$

## Interpretation of $U$ and $V$

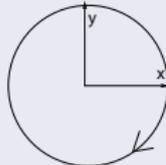
Linear +45



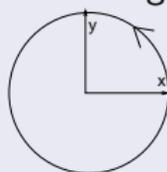
Linear -45



Circular Left



Circular Right

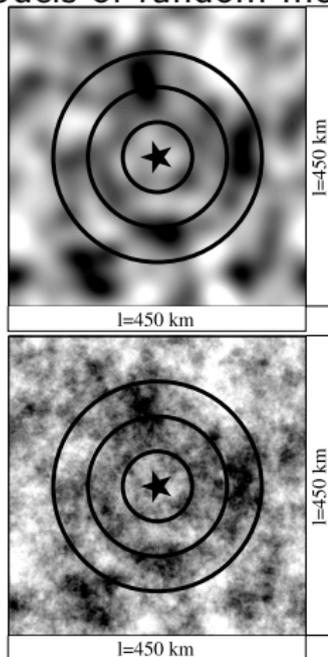


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# Numerical Tests

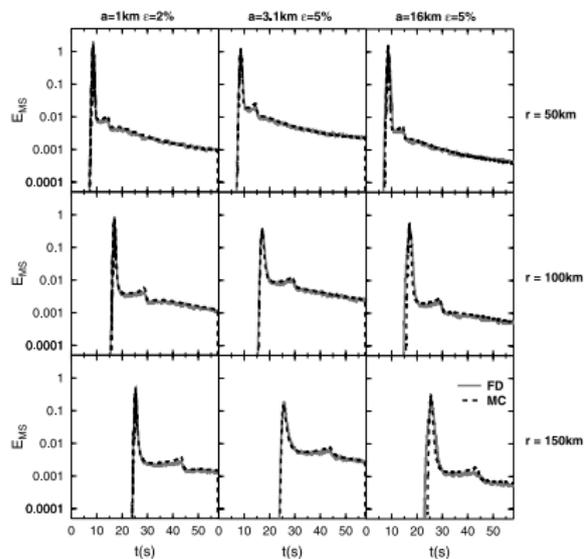
## Models of random media



- 2-D Gaussian, exponential media
- Finite-Difference calculations
- $P$  wave source
- Ensemble average
- Elastic R.T.E.
- $\mathbf{p}(\hat{\mathbf{k}}, \hat{\mathbf{k}}')$  from Born approx.
- Monte-Carlo simulations

*Przybilla et al.*, J. Geophys. Res, 111,  
B04305, 2006

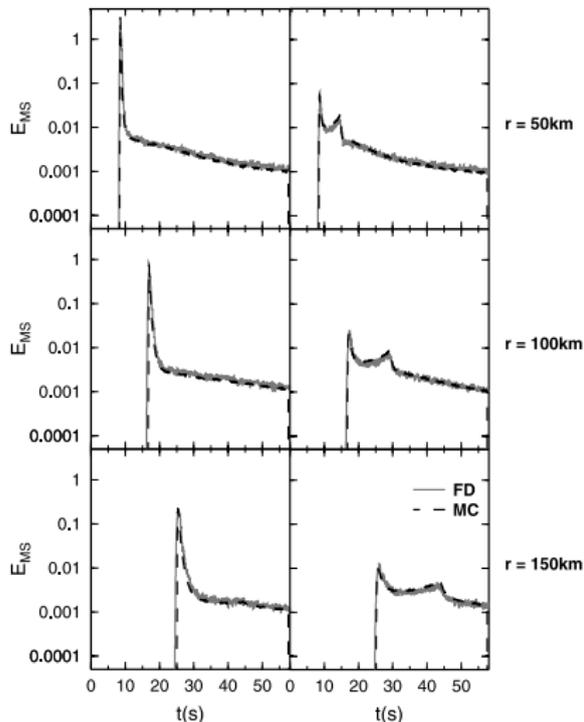
## Comparison between RT and FD simulations



- Left:  $t^S = 11.4s$ , weak anisotropy, weak scattering
- Middle:  $t^S = 0.5s$
- Right:  $t^S = 0.1s$ , strong anisotropy, strong scattering
- Excellent agreement even for strong forward scattering
- Uniform energy distribution at long time

Przybilla et al., J. Geophys. Res, 111, B04305, 2006

# Comparison between RT and FD simulations



- Excellent prediction of energy partition on radial and transverse components
- **Equipartition** at large time
- Energy ratios stabilize:

$$\frac{E_s}{E_p} = \left( \frac{v_p}{v_s} \right)^2$$

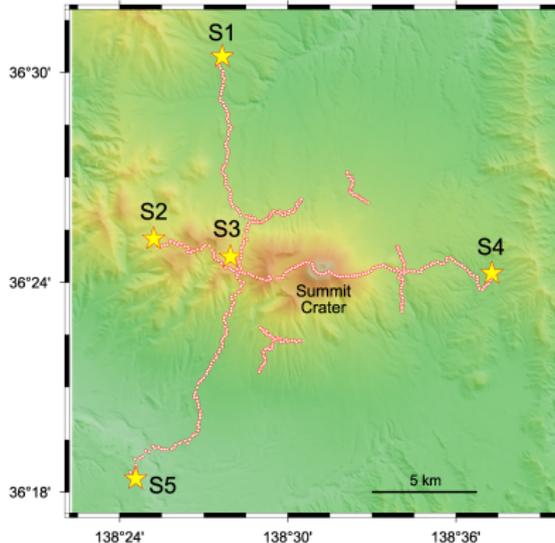
Przybilla *et al.*, J. Geophys. Res, 111, B04305, 2006

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# Imaging of volcanic heterogeneity

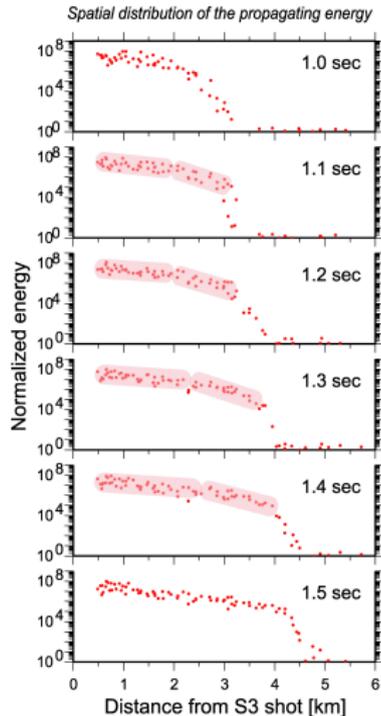
## Experiment at Osama Volcano (Japan)



- 5 Dynamite shots
- Vertical sensors 2Hz
- Station spacing: 50-150m
- Data normalized with late coda
- Shot 3 in the 8-16 Hz frequency band

Courtesy of M. Yamamoto

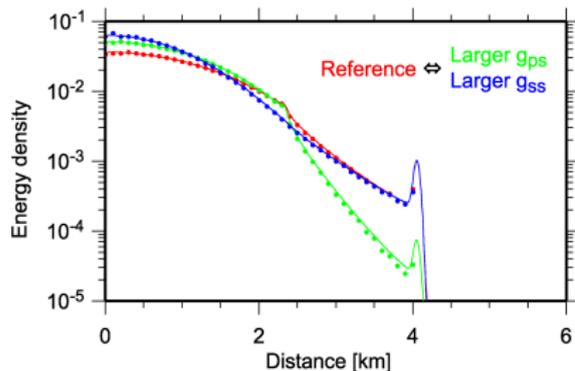
# Spatial Distribution of Energy



Courtesy of M. Yamamoto

- Observe the two slopes
- Shear waves are mode converted
- Spatial homogenization of energy at late time

# Role of the scattering parameters



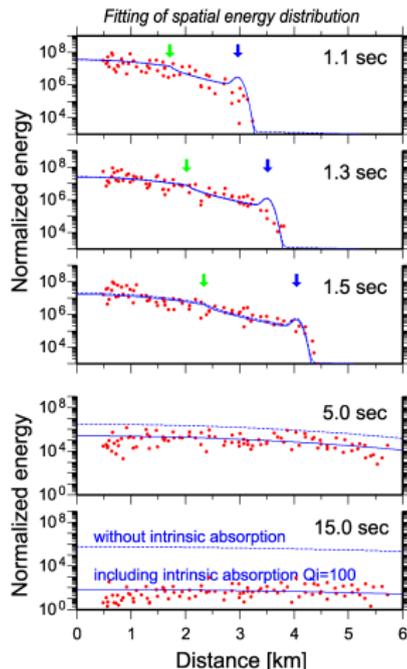
- **Reference:**  
 $v^P = 2.7 \text{ km/s}$ ,  $v^P/v^S = \sqrt{3}$ ,  $I^{PP} = 3$   
 $\text{km}$ ,  $I^{PP}/I^{PS} = 2$ ,  $I^{PP}/I^{SS} = 3$
- **Smaller  $I^{PS}$ :**  $I^{PP}/I^{PS} = 4$
- **Smaller  $I^{SS}$ :**  $I^{PP}/I^{SS} = 6$

- 3-D elastic radiative transfer
- Isotropic scattering
- Explosion: isotropic  $P$  source
- Reciprocity:  

$$I^{SP} = 2 \left( \frac{v^P}{v^S} \right)^2 I^{PS}$$
- Vary the ratios  $I^{PP}/I^{PS}$  and  $I^{PP}/I^{SS}$

Courtesy of M. Yamamoto

# Constraint on medium heterogeneity



Impressive agreement!

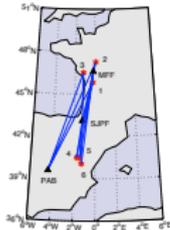
Yamamoto and Sato, to appear in J.G.R.

## Inversion results

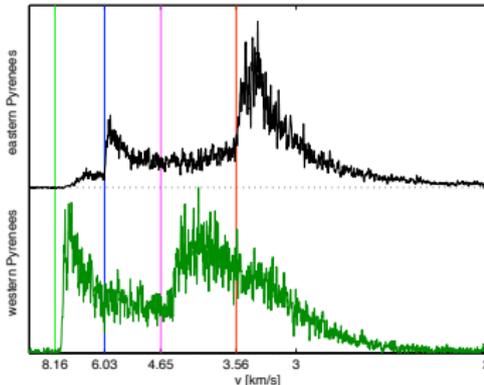
- $v^P = 2.7 \text{ km/s}$
- $l^{PP} = 3.8 \text{ km}$ ,  $l^{PP} = 2.4 l^{PS}$ ,  
 $l^{PP} = 3 l^{SS}$
- Absorption:  $Q_i = 100$
- Extremely strong scattering:  
 $l^P \approx l^S \approx 1 \text{ km}$
- Coupling parameters: rich information on the nature of heterogeneity

# Modeling of $L_g$ blockage through the Pyrenees

## Wavepaths through the Pyrenees

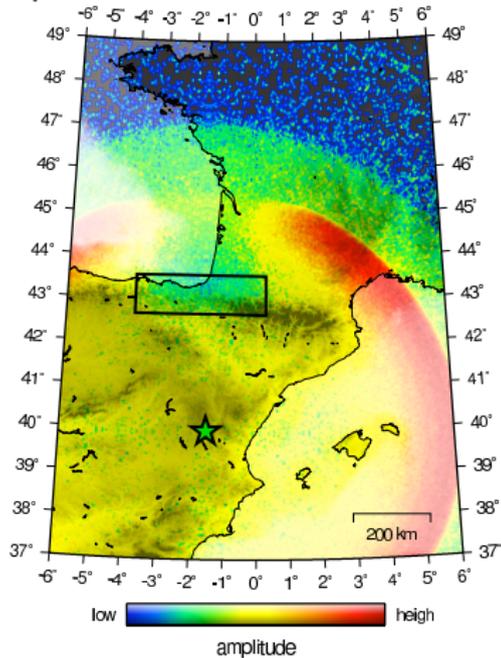


## Extinction of $L_g$ waves

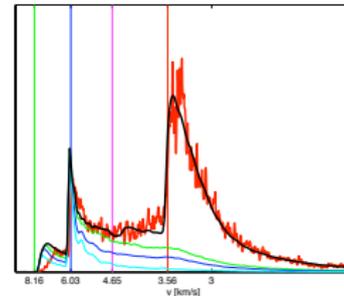


- Blockage in a localized zone
- Not explained by Moho jump or faults
- Monte Carlo simulations of RT
- 3-D anisotropic scattering and mode conversion
- Depth varying velocity and scattering properties

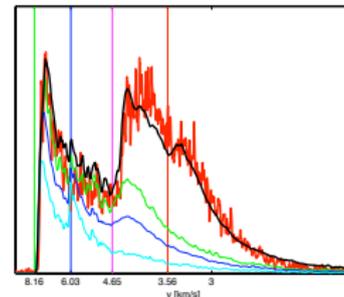
## Snapshot of numerical simulation



## Undisturbed zone



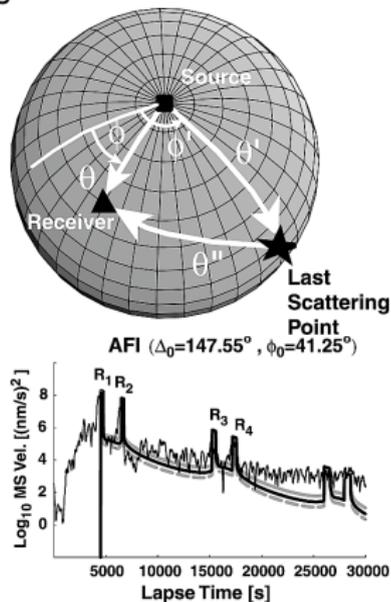
## Through anomaly



Sens-Schönfelder et al., J. Geophys. Res., 114, B07309, 2009

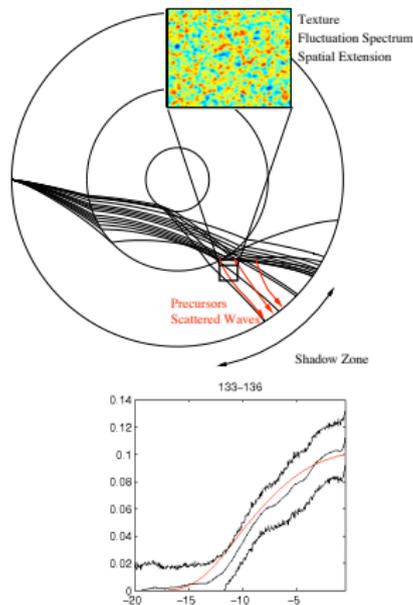
# Applications at the global scale

## Multiple scattering of Rayleigh waves



Sato and Nishino, J. Geophys. Res., 107, 2343, 2001

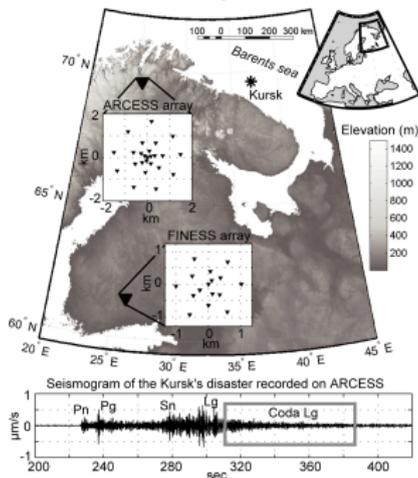
## Scattering in the lower mantle



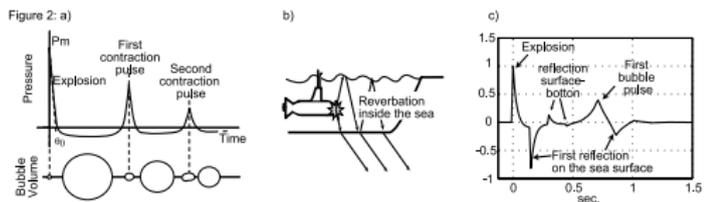
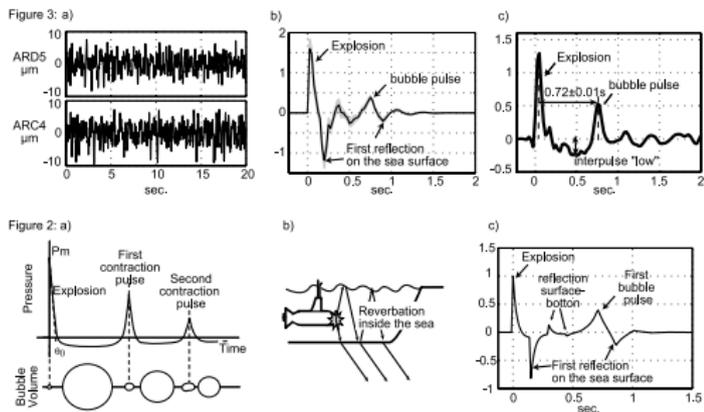
Margerin and Nolet, J. Geophys. Res., 108, 2514, 2003

# Application to source studies

## The Kursk Explosion



## Comparison between observed and "theoretical" explosion



Sèbe et al., Geophys. Res. Lett.,  
 32, L14308, 2005

## More radiative transfer

### Asymptotic solution of Bethe-Salpeter equation

Wigner distribution of two Green functions:

$$\Gamma(\mathbf{x}, \mathbf{r}; t, \tau) = \langle G(\mathbf{x} + \mathbf{r}/2, t + \tau/2) G(\mathbf{x} - \mathbf{r}/2, t - \tau/2)^* \rangle$$

Fourier transform over  $\tau$ :

$$C(\mathbf{x}, \mathbf{r}; t, \omega) = \int \Gamma(\mathbf{x}, \mathbf{r}; t, \tau) e^{-i\omega\tau} d\omega$$

Asymptotic result  $t \rightarrow \infty$

$$C(\mathbf{x}, \mathbf{r}; t, \omega) \sim \frac{e^{-\omega t/Q_i}}{(Dt)^{3/2}} \text{Im} \langle G(\mathbf{r}, \omega) \rangle$$

## Consequence 1: Equipartition

### Full-Space Elastic Green Function

Spectral domain:

$$\text{Im} \hat{G}_{ij}(\mathbf{k}, \omega) \sim (\delta_{ij} - \hat{k}_i \hat{k}_j) \delta(\omega^2 - v_s^2 k^2) + \hat{k}_i \hat{k}_j \delta(\omega^2 - v_p^2 k^2)$$

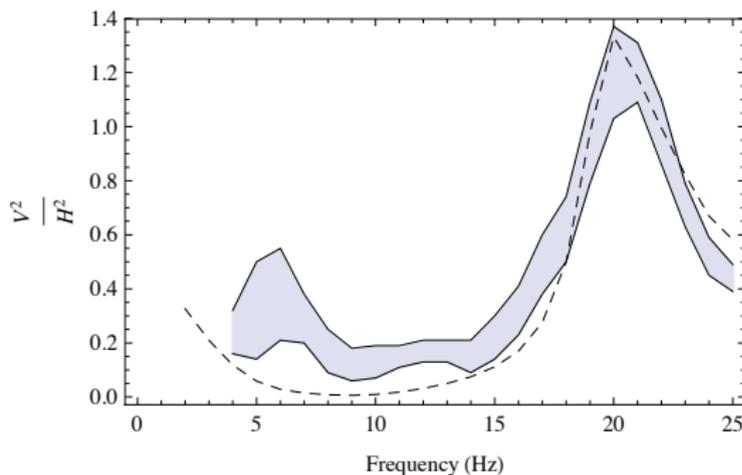
Total Energy:

$$\begin{aligned} \text{Trace} [\text{Im} G_{ik}(0, \omega)] &= \text{Trace} \left[ \iiint \text{Im} \hat{G}_{ij}(\mathbf{k}, \omega) d^3 k \right] \\ &\sim \text{Shear} + \text{Long} \\ &\sim \underbrace{2 \frac{\omega^2}{v_s^3}} + \underbrace{\frac{\omega^2}{v_p^3}} \end{aligned}$$

## Application of Equipartition to Site Effect Studies

Velocity Model
$h_1 = 4\text{m}$ $\alpha_1 = 300\text{m/s}$ $\beta_1 = 150\text{m/s}$ $\rho_1 = 2200\text{kg/m}^3$
$h_2 = 11\text{m}$ $\alpha_2 = 900\text{m/s}$ $\beta_2 = 500\text{m/s}$ $\rho_2 = 2200\text{kg/m}^3$
$h_3 = 50\text{m}$ $\alpha_3 = 3100\text{m/s}$ $\beta_3 = 1600\text{m/s}$ $\rho_3 = 2700\text{kg/m}^3$
$\alpha_\infty = 5400\text{m/s}$ $\beta_\infty = 3000\text{m/s}$ $\rho_\infty = 2700\text{kg/m}^3$

### Energy partition in the coda

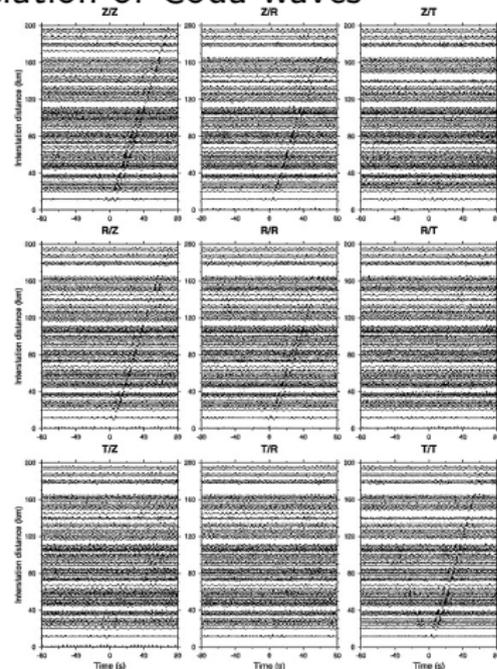
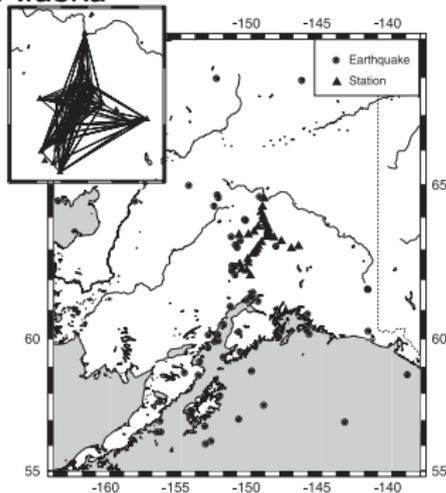


Margerin et al., Geophys. J. Int., 177, 571-585, 2009

## Consequence 2: Green function reconstruction

### Correlation of Coda waves

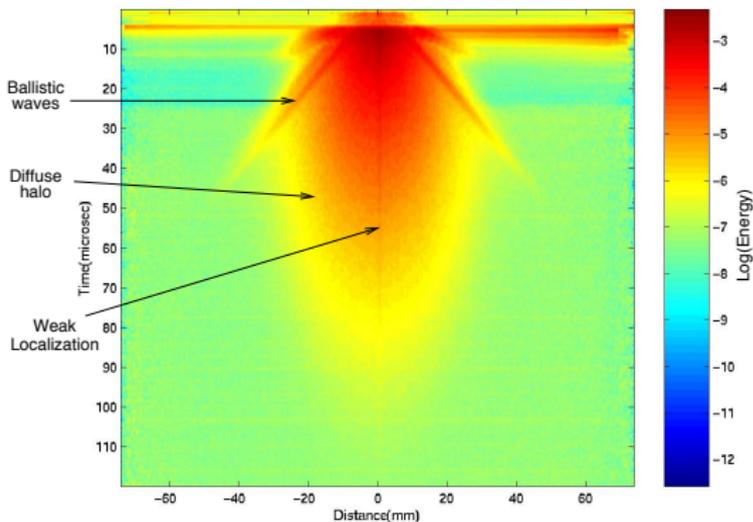
#### Temporary experiment in Alaska



Paul et al., J. Geophys. Res., 110, B08302, 2005

# Beyond radiative transfer: Weak localization

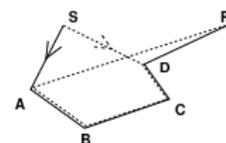
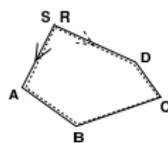
## Ultrasound propagation in sand



Courtesy A. Derode, L.O.A.

## Interference of reciprocal paths

Configuration 1



Configuration 2

